

Analysis of Convection Initiation Processes in Complex Terrain with the Synergy of COPS Remote Sensing Data

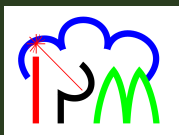
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Rhine Valley

Black Forest Mountains

COPS Supersite H



Initiation of (Deep) Convection (CI)

Analyses ahead of COPS indicated that CI in complex terrain is mainly determined by

- the convergence and updrafts created by forced lifting on the windward side and thermally-forced anabatic flow,
- the wind shear profile in the region of the ridges,
- variations in the depth of the convective boundary layer as well as in moisture, convective inhibition (CIN), and convective available potential energy (CAPE) across the mountain ridges, ←
- the presence of gravity waves impinging on the ridges,
- aerosol loading in the pre-convective environment influencing the diurnal cycle of boundary variables.

How is CI handled by models? Cumulus Parametrization!

CI trigger function in mass flux convection schemes tests each grid point by

- vertical velocity offset at LCL (= cloud base), Kuo 1965
- adding temperature offset depending on grid-scale vertical wind (large scale lifting supports CI), Fritsch and Chappel 1980
- height depending threshold of temp. offset, Kain 2003

- shallow/deep convection depending on range between LCL and LNB of the perturbed case (=extend of subgrid-scale cloud) precip. is turned on; typically 3 km

Presently used CI trigger functions are justified empirically!

Many models trust on just a single parameter (e.g., COSMO with Tiedke 1989 uses offset to vertical wind only).



What do we need to investigate CI?

1. Accurate data, especially of water vapor (lesson from, e.g., IHOP_2002)

→ Intercomparisons, higher-order corrections to reach better than 5 % accuracy

2. IOP case studies

→ Investigate small scale heterogeneity of water vapor, temperature, wind, fluxes, boundary layer height, clouds, aerosols and their relation to CI

→ Combine simultaneous data of temperature and water vapor:

$d\theta/dz$, $d\theta_v/dz$, buoyancy, CAPE, CIN

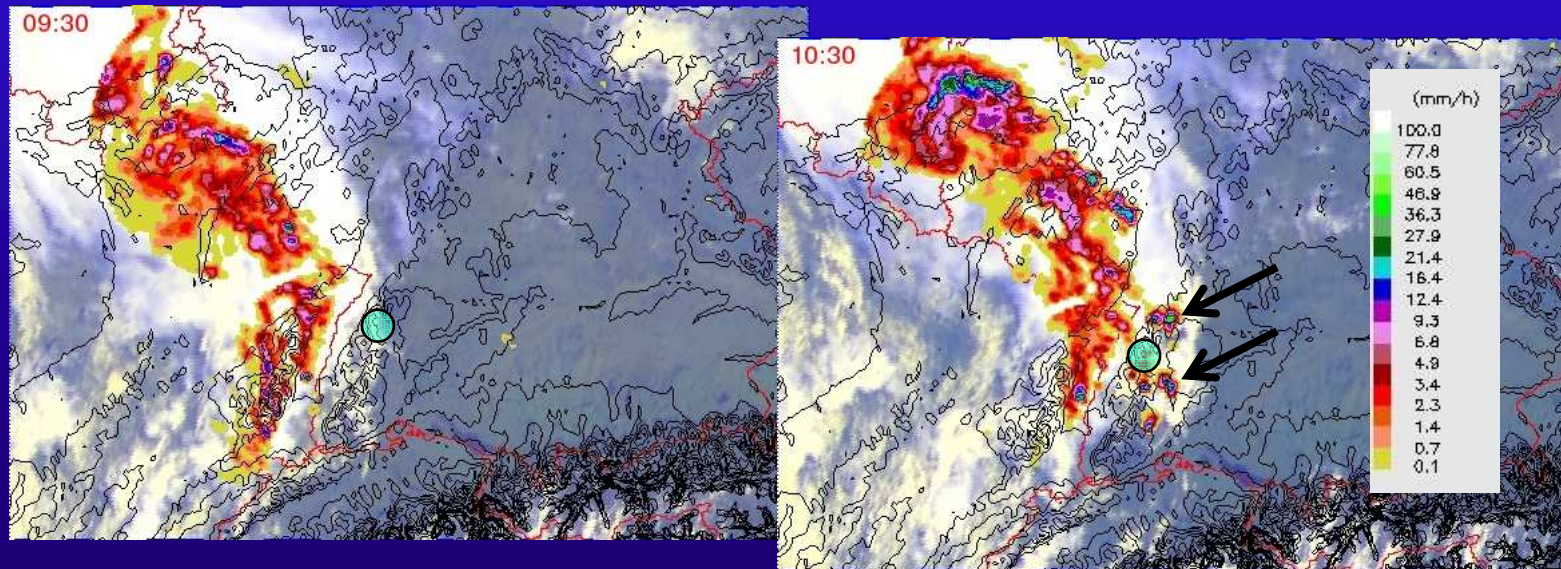
3. Comparison with different parameterization concepts

→ D-PHASE, COPS-GRID re-analyses, and hybrid convection schemes

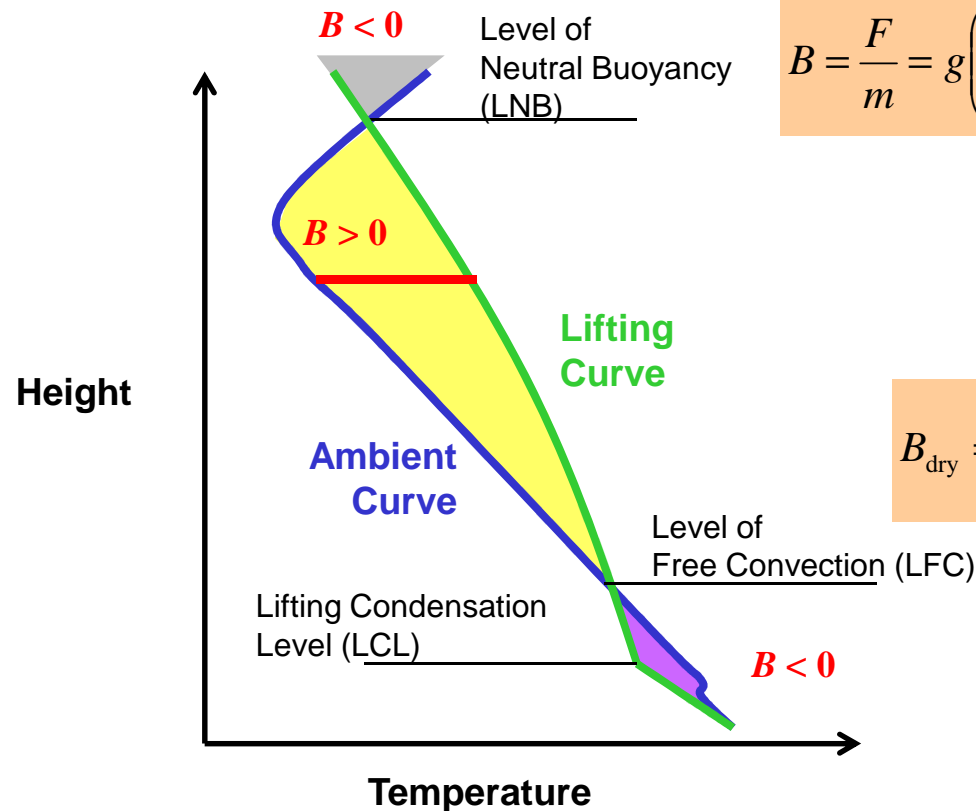


IOP 9c, 20 July 2009

- COPS IOP 9c: flooding, modification of front in Northern Black Forest, all COPS research instruments operated, „non-case“ on Hornisgrinde
-> perfect conditions to measure pre-convective fields with lidars
- New product: 5-minute Buoyancy profiles of collocated rotational Raman lidar (T) and water vapor DIAL (& ground met station)
- Precision (stat. uncertainty) directly obtained from the signal intensities
- Differences to drifting radiosondes?



Buoyancy B



$$B = \frac{F}{m} = g \left(\frac{\rho_{\text{parcel}} - \rho_{\text{ambient}}}{\rho_{\text{ambient}}} \right) = g \left(\frac{\theta_{\text{virtual,parcel}} - \theta_{\text{virtual,ambient}}}{\theta_{\text{virtual,ambient}}} \right)$$

$$\theta_{\text{virtual}} = T \left(\frac{1000 \text{ hPa}}{p} \right)^{0.286} (1 + 0.61 r)$$

$$B_{\text{dry}} = g \left(\frac{\rho_{\text{parcel,dry}} - \rho_{\text{ambient,dry}}}{\rho_{\text{ambient,dry}}} \right) = g \left(\frac{\theta_{\text{parcel}} - \theta_{\text{ambient}}}{\theta_{\text{ambient}}} \right)$$

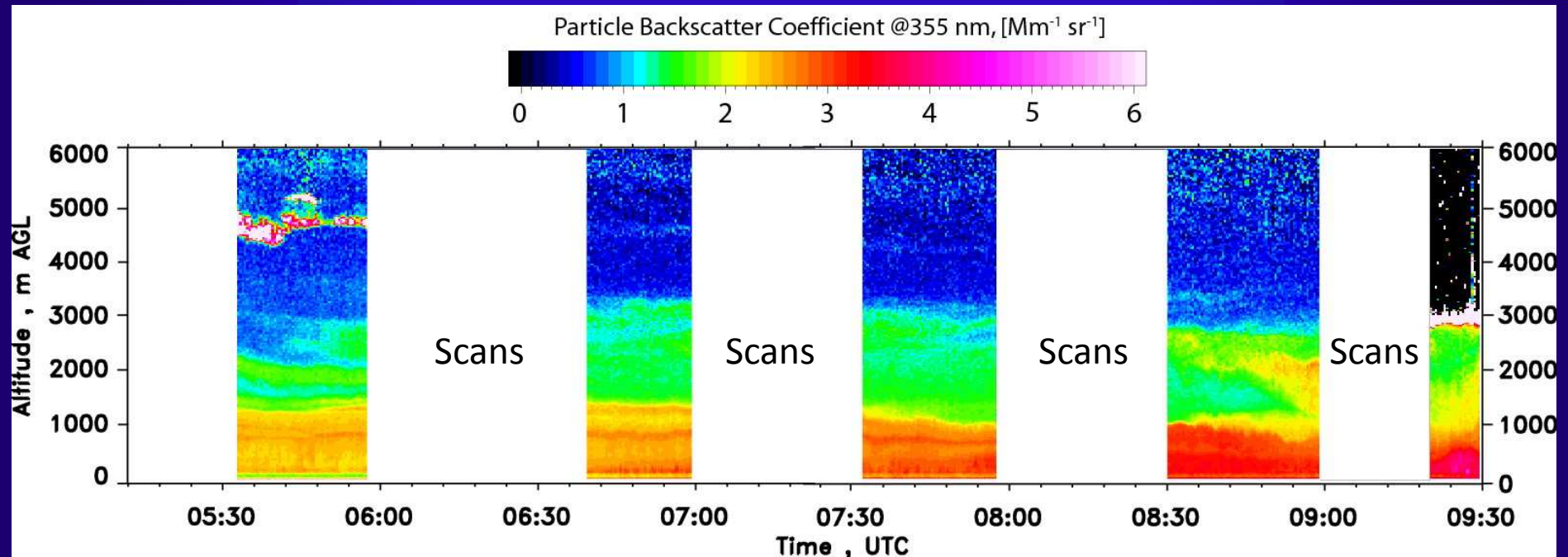
$$\theta = T \left(\frac{1000 \text{ hPa}}{p} \right)^{0.286}$$

Ambient Curve: T and r from lidar

Lifting Curve: T and r from ground met station

Particles: UHOH Rotational Raman Lidar

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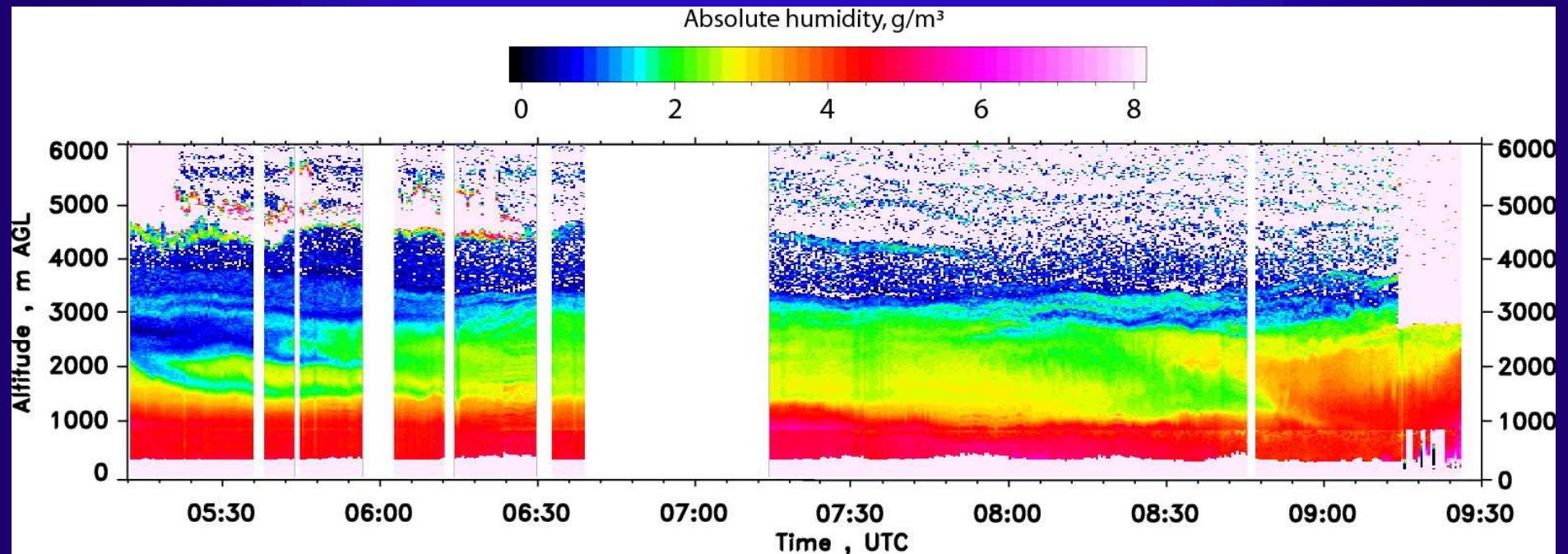
$\Delta z = 3.75$ m

$\Delta t = 13$ s



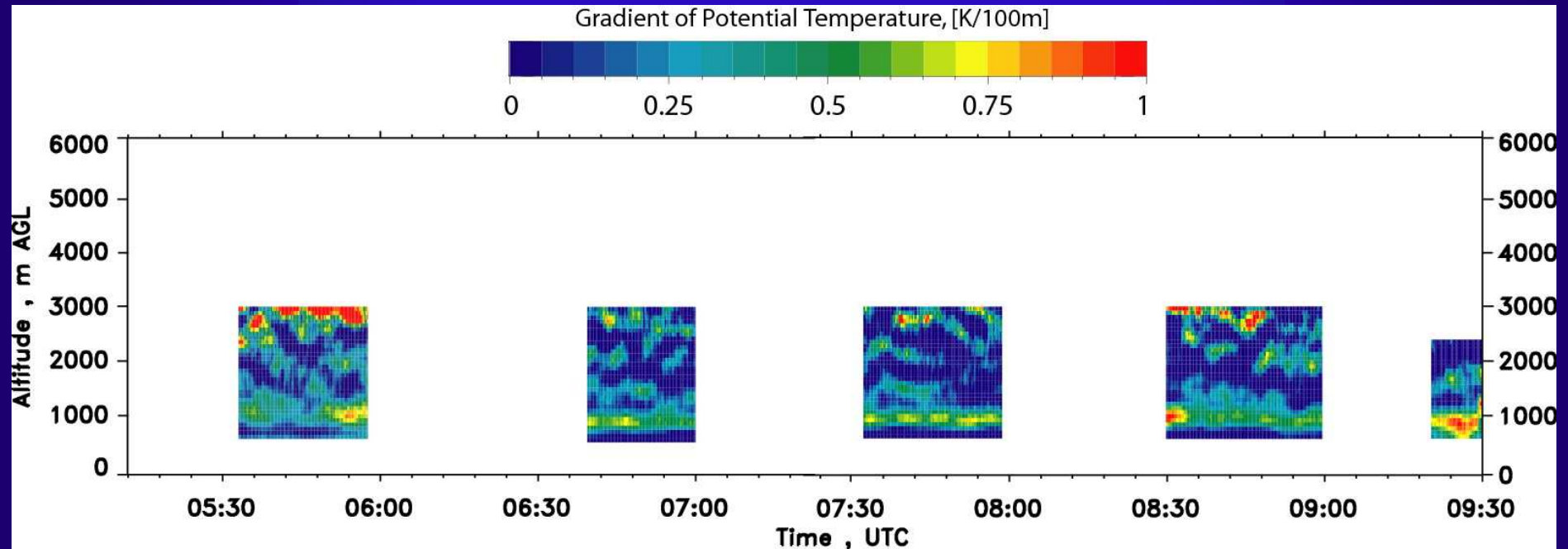
Humidity: UHOH DIAL

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$d\theta/dz$: UHOH Rotational Raman Lidar

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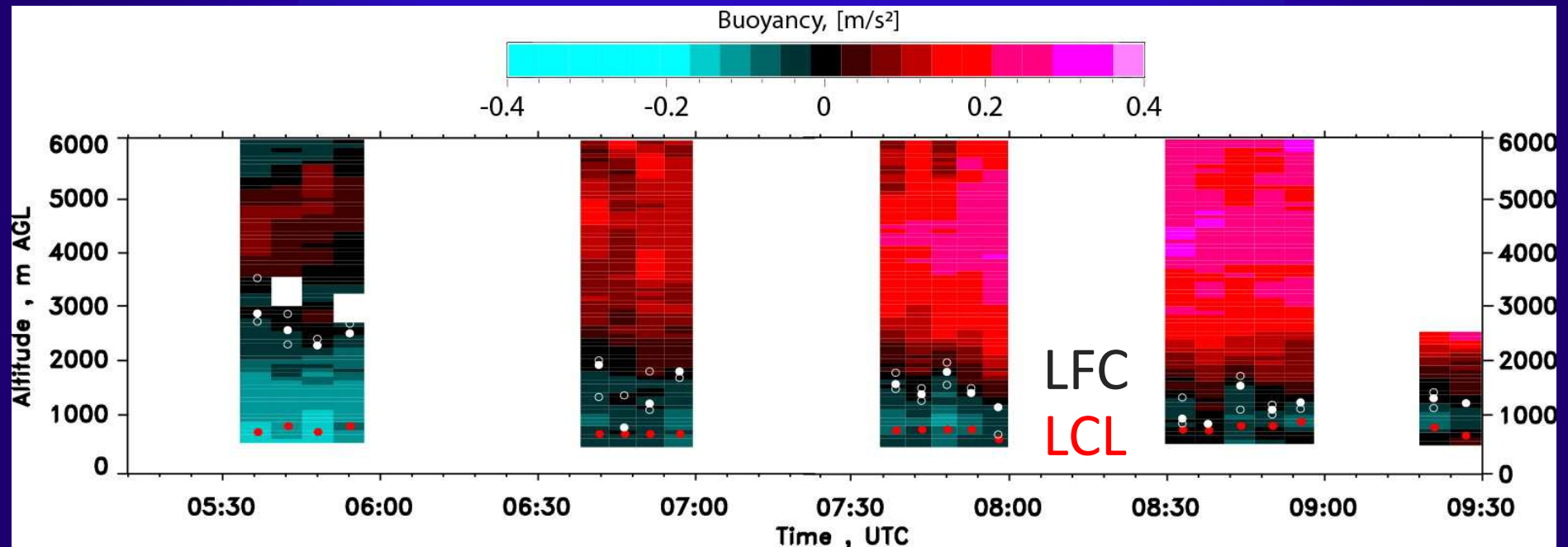
$$\theta = T \left(\frac{1000 \text{ hPa}}{p} \right)^{0.286}$$

$\Delta z = 75 \text{ m}$
 $\Delta t = 3 \text{ min}$



„Dry“ Buoyancy: UHOH Rotational Raman Lidar

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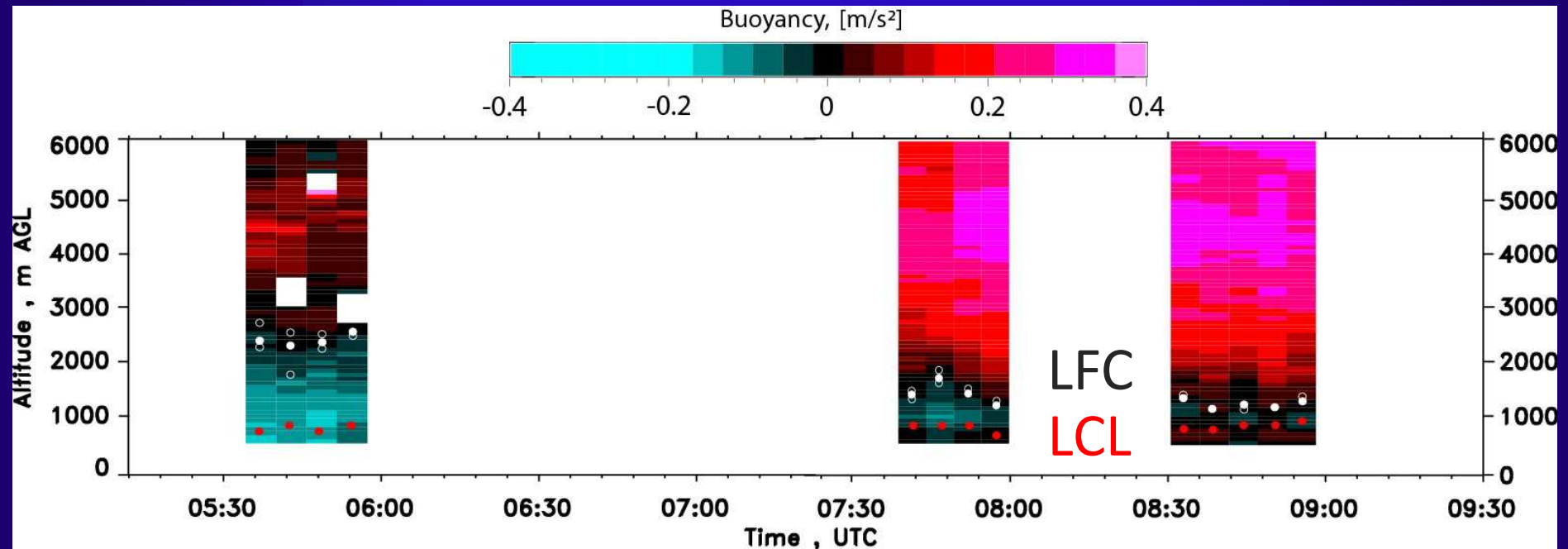
$$B_{\text{dry}} = g \left(\frac{\theta_{\text{parcel}} - \theta_{\text{ambient}}}{\theta_{\text{ambient}}} \right)$$

$\Delta z = 150 \text{ m}$
 $\Delta t = 5 \text{ min}$



„Moist“ Buoyancy: UHOH RRL & DIAL

20 July 2007, IOP 9c



$$B = g \left(\frac{\theta_{\text{virtual,parcel}} - \theta_{\text{virtual,ambient}}}{\theta_{\text{virtual,ambient}}} \right)$$

$\Delta z = 150 \text{ m}$
 $\Delta t = 5 \text{ min}$



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Lidar ✓
Radiosondes (✓)

Bhawar, Di Girolamo et al., Tue, 12:45
Bhawar, Di Girolamo et al., poster

2. IOP case studies

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→ Combine simultaneous data of temperature and water vapor:

$d\theta/dz$, $d\theta_v/dz$, buoyancy, CAPE, CIN

Ongoing at
several
groups

Ongoing

3. Comparison with different parameterization concepts

→ D-PHASE, COPS-GRID re-analyses, and hybrid convection schemes

Outlook.

